ISOPACH MAP OF QUATERNARY AND UPPER TERTIARY STRATA

into sopacim map of quaternary and upper lettary strate retrievathe near-surface stratigraphy and structure of the Outer Continental Shelf (OCS) Oil and Gas Lease Sale 57 area. The seismic systems used in this interpretation had either a minisleeve exploder or a 15-cubic-inct watergun as the energy source. The incoming signal was sampled at a 0.5-ms rate for 1 s. The data were processed using a twelve-fold, common-depth-point (CDP) stack, and were displayed in both automatic and nootrol and relative true annibuse formats:

Strationanhy

Morton Basin is filled with about 4 km of sediment as old as earl; Iertiary (Melson, Hopkins and Scholl, 1974) or Late Cretaceous (Fishe and others, 1979). These sediments are continuous with the main layers sequence (Scholl and Hopkins, 1969) and overlie the acoustic basement or Paleozoic rocks. On the basis of velocity contrasts identified fro sombobuy refraction surveys, Fisher and others (1979) have divided this correspond to various lithologic units that were deposited as the basis of velocity of the youngest velocity unit, A, is characterized by compressional velocities of 1.7 to 2.1 km/s, and the underlying unit, B by velocities of 2.3 to 2.8 km/s. They do not specifically explain th A and B velocity contrast, nor describe the lithology of unit A However, they imply that the older unit, B, was deposited during marine transgression of the basin. The unconformity at the base of unit B marks the time when the basin changed from a nommarine or deltain B marks the time when the basin changed from a nommarine or deltain Scholl and Hopkins (1969) and Holmes and others (1979), and as late Mocene boshol with the second process of unit A is probably younger than the Oligocene-Mocene boundary and may be younger than the Mocene.

corresponding to unit in the survey area. We have tentative assigned an age of Quaternary and late Tertiary to the strata. We located the base of unit A by using the A and B velocity contrast, and corresponding change in reflective characteristics. On the seism sections, the base of unit A is displayed as a narrow zone ocntinuous, high-amplitude reflectors which occurs at depths rangin from 180 to 440 m. This zone of reflectors was not seen over much o the survey area because of the masting effect of extensive, shallow acoustic anomalies (Steffy and Hoose, 1981). However, we infer the Quaternary and upper Tertiary strata, unit A, are characteristical continuous, parallel, high-amplitude features, whereas reflectors belo these strata are discontinuous, low-amplitude features. A subtil divergence between the internal reflectors of units A and B towards the basin axis has been observed. This has also been recognized by Schol (oral communication, 1981), using seismic data collected by the Geologi Diviston, U.S. Geological Survey.

An acoustic velocity of 1,700 m/s for the strata of unit A was used to convert from travel time to isospach thickness. This velocity was obtained from the multichannel velocity analysis and matches the sonobuoy refraction results of Holmes and others [1979]. Holmes and others propose that this velocity indicates loosely cemented sands, silts, and clays deposited in a marine environment. This assumption is corroborated by two boreholes drilled by the U.S. Bureau of Hines in 1867 (Scholl and Noptins, 1869). These holes, drilled to depths of 61.0 and 73.2 m (200 and 200 feet), are penetrated it.2 and 22.9 m (40 and 75 feet) of Quaternary sediments underlain by marine silts and clays dated by pollen, foraminifera, and ostracods as Pilocene.

commonly occur along the base of unit A which ranges in depth between 100 and 260 as (Steffy and Hoose, 1981). Acquistic anomalies also occur thin the upper 100 as two-way travel time (140 m) of the seismit section. Holmes and Thor (1981), and Steffy and Hoose (1981) have explained that many of these shallow anomalies result from gas-charge sediments or a strong reflecting horizon. The gas-charged sediments and the horizon correspond to a buried, Quaternary peat layer which has bee sampled in cores.

We attribute the stratigraphic relationship and the velocity contrast between units A and B to a change in the upper Tertiary sedimentary environment of the Morton Basin. This interpretation is sedimentary environment of the Morton Basin. This interpretation is intermed. The sedimentary of the two units; (2) the nearly parallel relationship between the zone of reflectors that define the A and I relationship between the zone of reflectors of unit 8; and (3) the inferred boundary and the internal reflectors of unit 8; and (3) the inferred presence of marine sediments in both units. This change in the sedimentary environment could have been caused by a change in the rate of subsidence according to Greene and Perry (Holmes and others, 1974, 97, 241) or a change in the sediments source (Nelson and others, 1974, 97, 241) or a change in the sediments source (Nelson and others, 1974, 1974).

Structu

The Quaternary and upper Tertiary strata show little structure deformation. These strata dip gently southward towards the east-weterending basin axis in the eastern portion of the map, and disouthwestward towards the northwest-southeast trending basin axis in the vestern portion of the map. Superimposed on these general dips at measural, elongated depocenters that parallel the basin axis. These smal depocenters are usually associated with growth faults that were mapped by Steffy and Hoose (1981) and are shown on this map. These faults havertical offsets of usually less than 50 m and are downthrown towarms the basin axis in the eastern portion of the map and are downthrown awfrom the basin axis in the extern portion of the map. Over most che area, these growth faults are parallel to deeper faults that do not the unit A strata. The deeper faults occur as sets forming graber or as single faults downthrown towards the basin axis.

REFERENCES CITE

Fisher, M. A., Patton, W. W. Jr., Thor, D. R., Holmes, M. L., Scott, E. W., Melson, C. H., and Wilson, C. L., 1979, Resourr report for proposed OCS Lease Sale 57: Norton Basin, Alaska: U.S. Geological Survey Open-File Report 79-720, 45 p.

Holmes, M. L., Cline, J., and Johnson, J. L., 1978, Geological setting of the Norton Basin gas seep: Proceedings of the 10th Annual Offshore Technology Conference, Houston, Texas, 1978, v. 1, p. 73-80.

Assessment of the Alaskan Continental Shelf, Annual Reports of Principal Investigators for the year ending March, 1980, v. 5, p. 261–285.

Seas: New York, Springer-Verlag, p. 119-140.

Scholl, D. W., 1981, Oral communication: Marine Geology Branch,
Geologic Division, U.S. Geological Survey, Menlo Park, CA.

Geologic Division, U.S. Geological Survey, Menlo Park, CA.

Steffy, D. A. and Hoose, P. J., 1981, Map showing acoustic anomalies and near-surface faulting, Norton Sound, Alaska: U.S. Geological

OPEN-FILE REPORT SERIES ON HORTON SOUND, ALASKA, 1981 **EXPLANATION** BERING SEA NORTON SOUND DATA OBSCURED BY SHALLOW STUART YUKON RIVER DELTA edited for conformity with Geological Survey editoral standards. Any use of trade names is for descriptive purposes only, and does not constitute endorsement of these products by the Geological Survey. UTM ZONE 3 MAP PROJECTION UTM CLARKE 1866 SPHEROID, ZONE3.

UTM ZONE 3

SOURCE OF SHORELINE FROM BLM PROTRACTION DIAGRAMS NQ3-7, NQ3-8, NP3-1 AND NP3-2. PUBLISHED IN 1976.

